

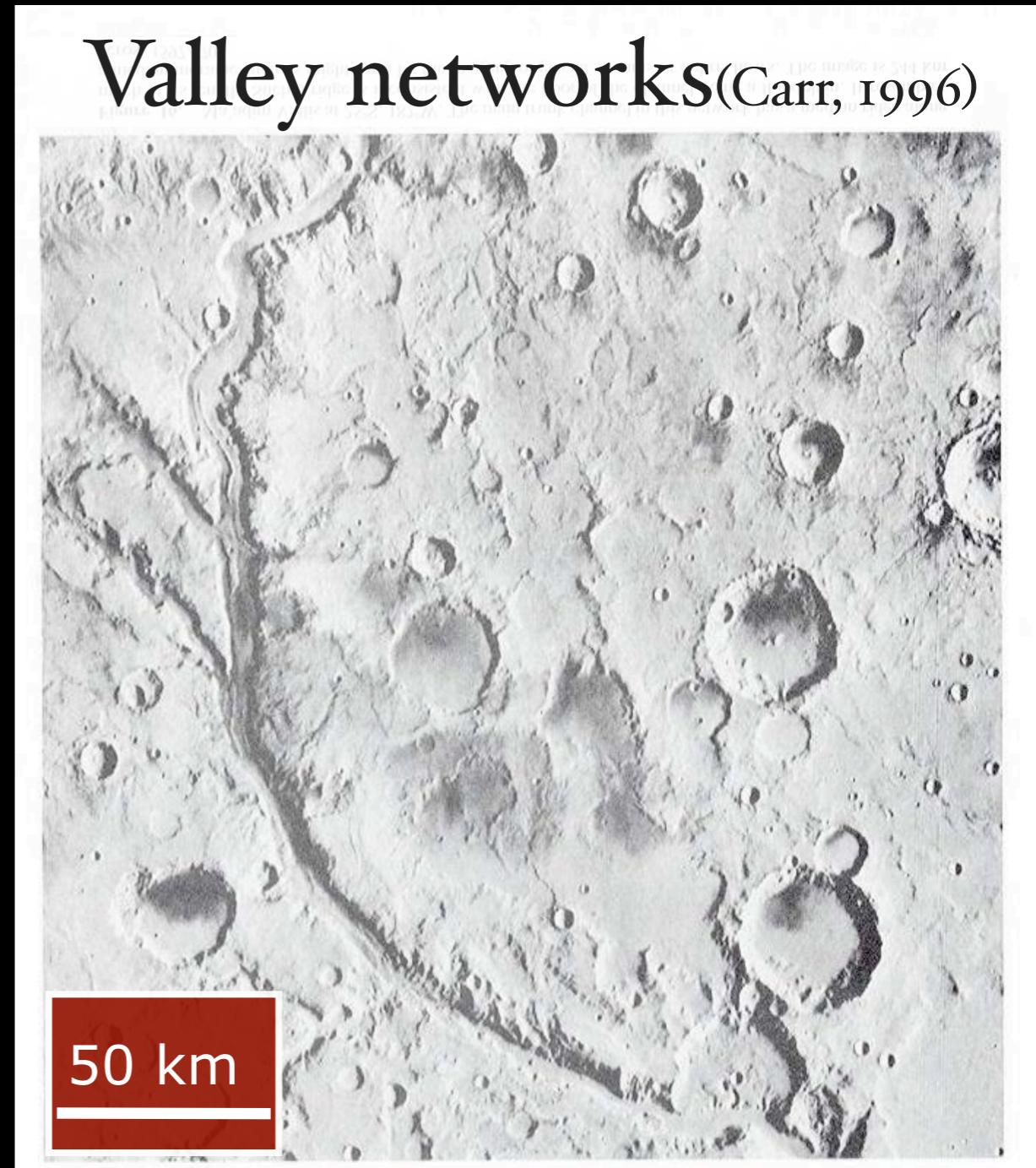
# **Warming of early Mars induced by CO<sub>2</sub> ice clouds**

Kobe University, the 21st century COE program  
Microsymposium, Climate and Meteorology on Mars  
Friday, June 17, 2005

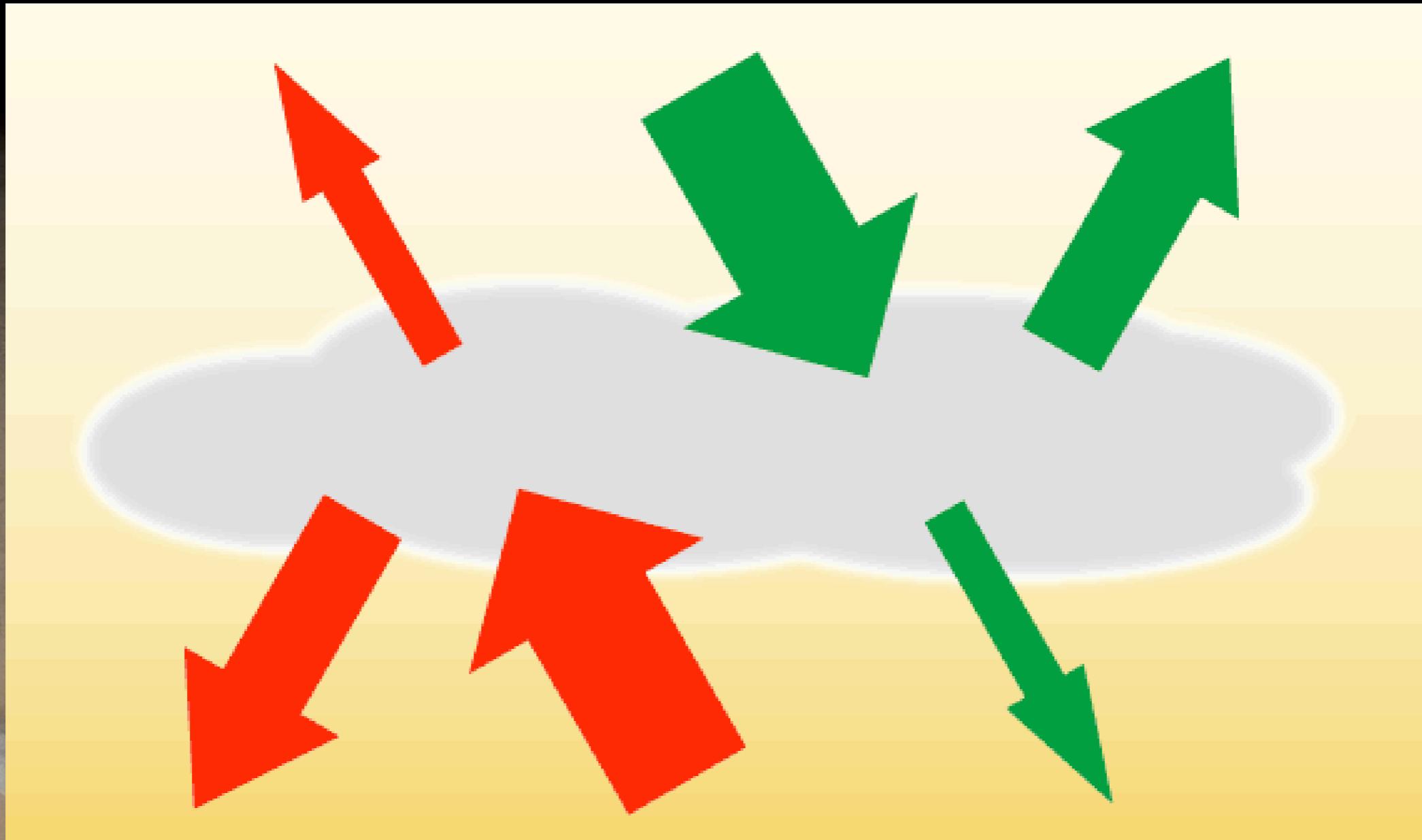
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# The faint young Sun paradox on Mars

- Early Mars ( -- 3.8 Ga)
  - wet and warm climate?
- The warm climate cannot be sustained (Kasting, 1991)
  - CO<sub>2</sub>-H<sub>2</sub>O atmosphere
  - Solar Luminosity 75 % as present
  - neglect radiation processes of cloud



# Scattering greenhouse effect of CO<sub>2</sub> ice cloud



cloud reflects IR radiation > Solar radiation

# Scattering greenhouse effect of CO<sub>2</sub> ice cloud

- Previous studies have shown
  - The strength of greenhouse effect depends on cloud parameters such as particle size and optical depth (e.g. Pierrehumbert and Erlick 1998)
  - Climate can become warm when cloud has optimal parameters (e.g. Mischna *et al.*, 2000)
  - Mechanisms determining these values have not been examined.

# mechanisms of cloud parameters change

- Cloud particle size and optical depth are changed by ...
  - particle coalescence by collision
  - particle evaporation as getting out of the cloud
  - particle growth by radiative cooling or evaporation by radiative heating in cloud layer

# mechanisms of cloud condition change

- Cloud particle size and optical depth are changed by ...
  - particle combination with collision
  - particle evaporation as getting out of the cloud
  - particle growth by radiative cooling or evaporation by radiative heating in cloud layer

This study:  
Estimation the greenhouse effect  
of stable CO<sub>2</sub> ice cloud

# 1-D radiation model

Solar luminosity:  
0.75 times of today  
(Gouth, 1981)

Two-stream approximation  
(cloud layer:  $\delta$ -Eddington approximation)

Albedo: 0.216  
(Kieffer et al. 1977)

CO<sub>2</sub>-H<sub>2</sub>O  
atmosphere

Cloud particle:

\* Mie theory

Complex indices of CO<sub>2</sub> ice (Warren, 1986)

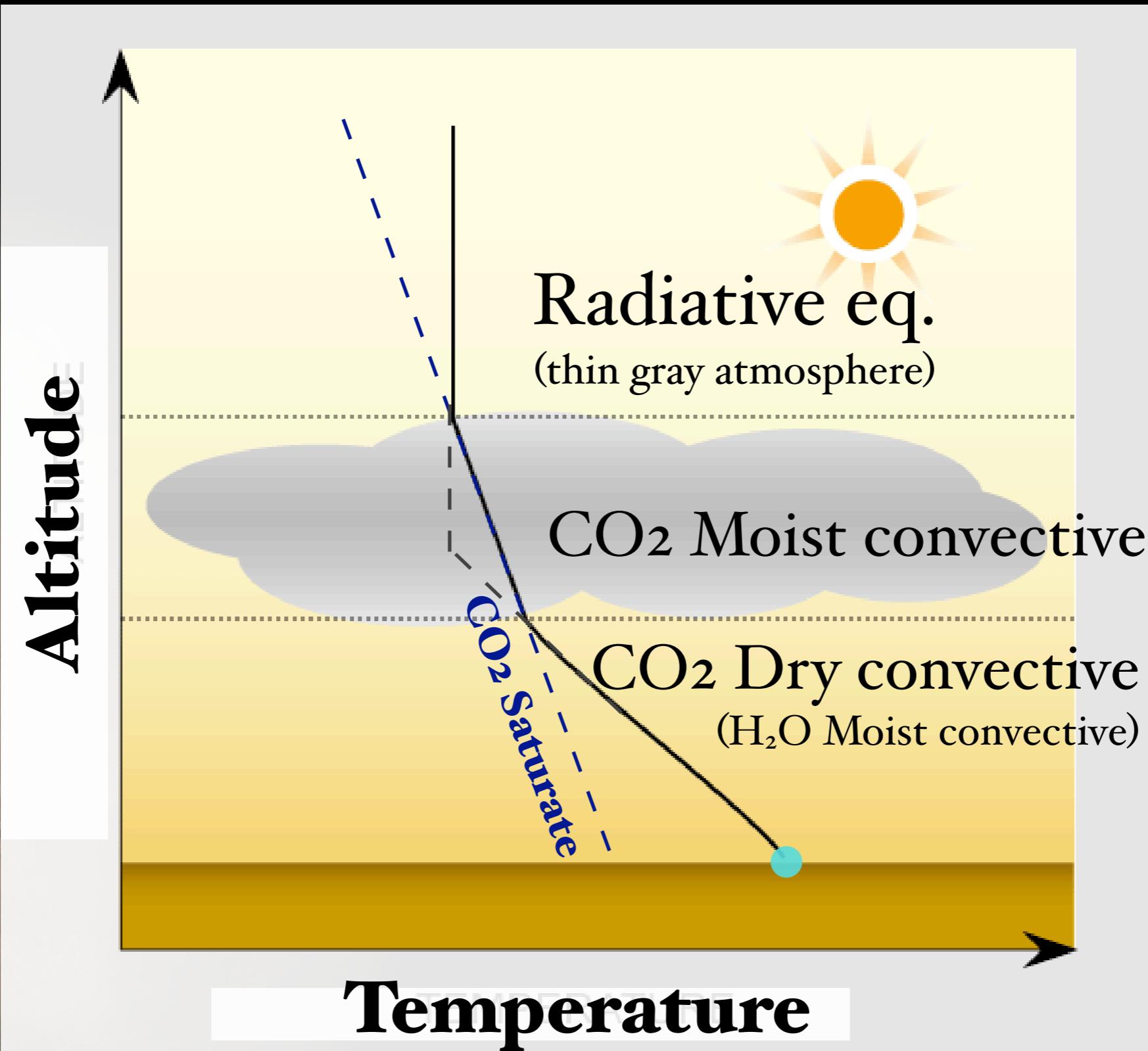
Gas ( for IR radiation only):

\* Line-by-line method

absorption line parameters (HITRAN<sub>2000</sub>)

\* At cloud layer, Random model  
band parameter (Houghton, 2002)

# Vertical temperature profile



# 1-D radiation model

Solar luminosity:  
75 % current value  
(Gouth, 1981)

Two-stream approximation  
(cloud layer:  $\delta$ -Eddington approximation)

**latent heat of CO<sub>2</sub> condensation  
= net cooling energy in cloud layer**

CO<sub>2</sub>-H<sub>2</sub>O  
atmosphere  
regolith Albedo: 0.216  
(Kieffer et al. 1977)

Cloud particle:

\* Mie theory

Complex indices of CO<sub>2</sub> ice (Warren, 1986)

Gas ( for IR radiation only):

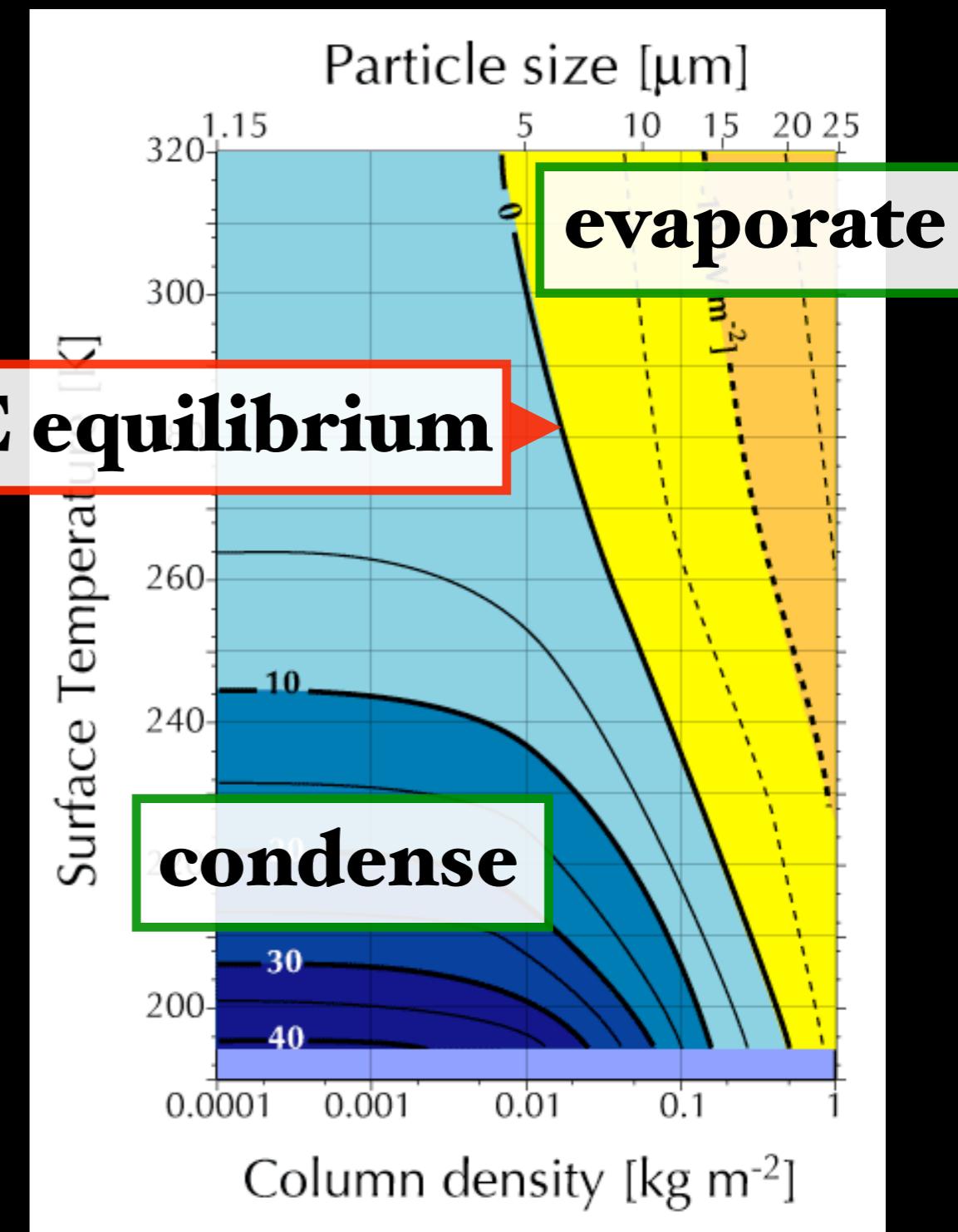
\* Line-by-line method  
absorption line parameters (HITRAN<sub>200</sub>)

\* At cloud layer, Random model  
band parameter (Houghton, 2002)

# Latent heat of CO<sub>2</sub> condensation

atmospheric pressure: 1 bar ; condensation nucleus:  $10^{10} \text{ m}^{-2}$

- Inverse correlation between particle size and latent heat is important
- By a negative feedback, the particle size change to the CE-equilibrium value if CO<sub>2</sub> condensate and evaporate quickly.



# Estimation of temperature

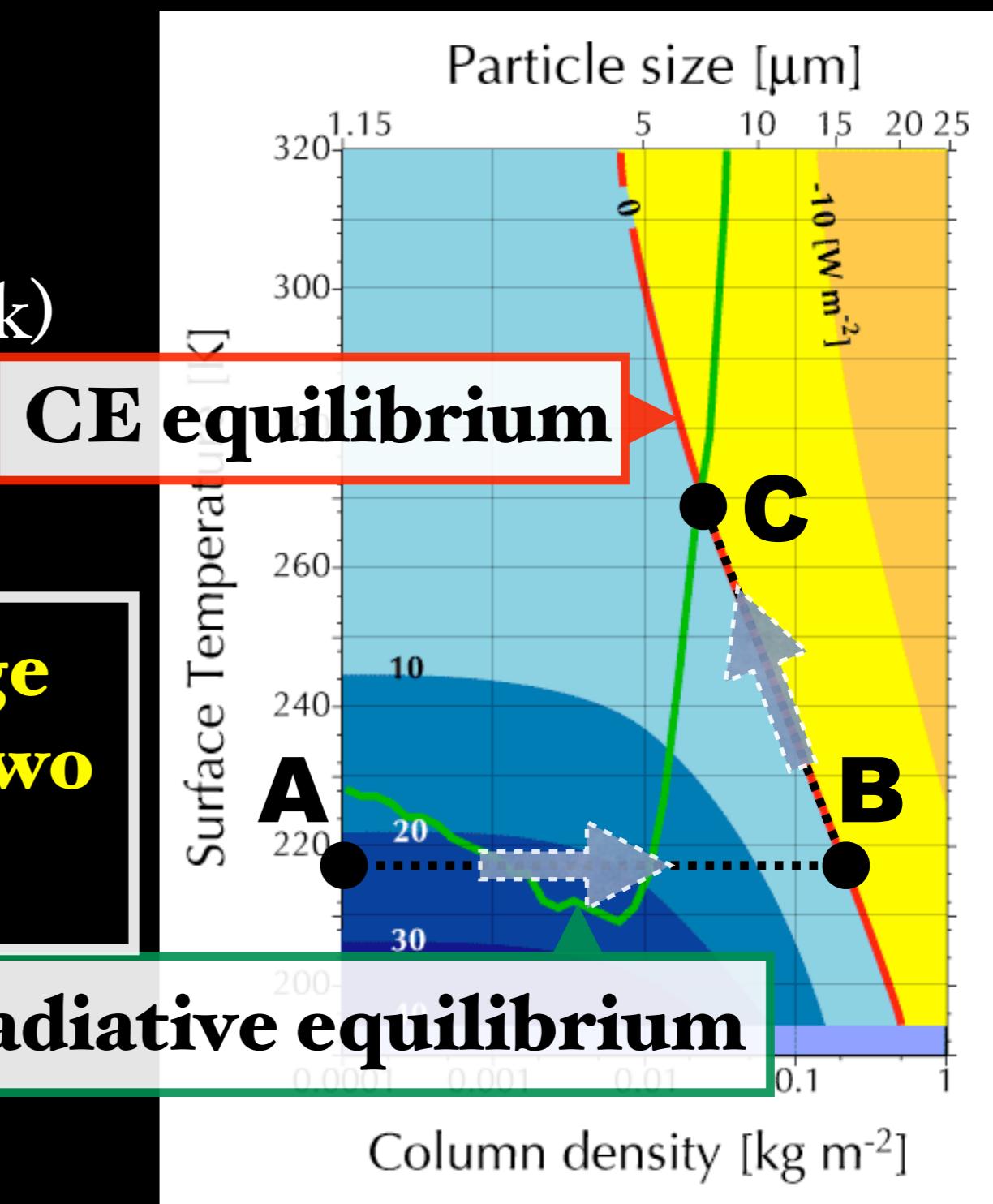
atmospheric pressure: 1 bar ; condensation nucleus:  $10^{10} \text{ m}^{-2}$

time scale

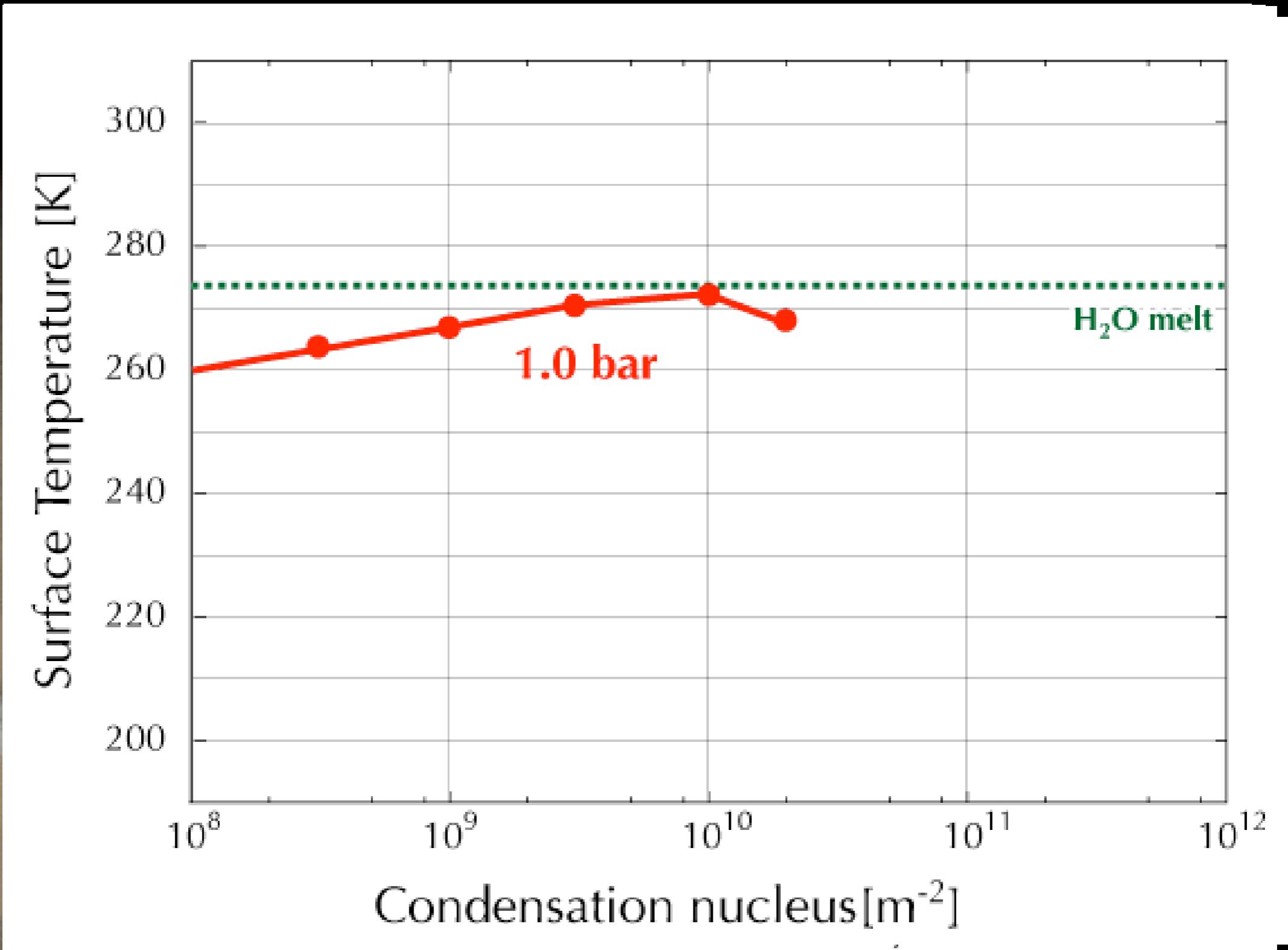
CE equilibrium(hour)

<< R equilibrium(week)

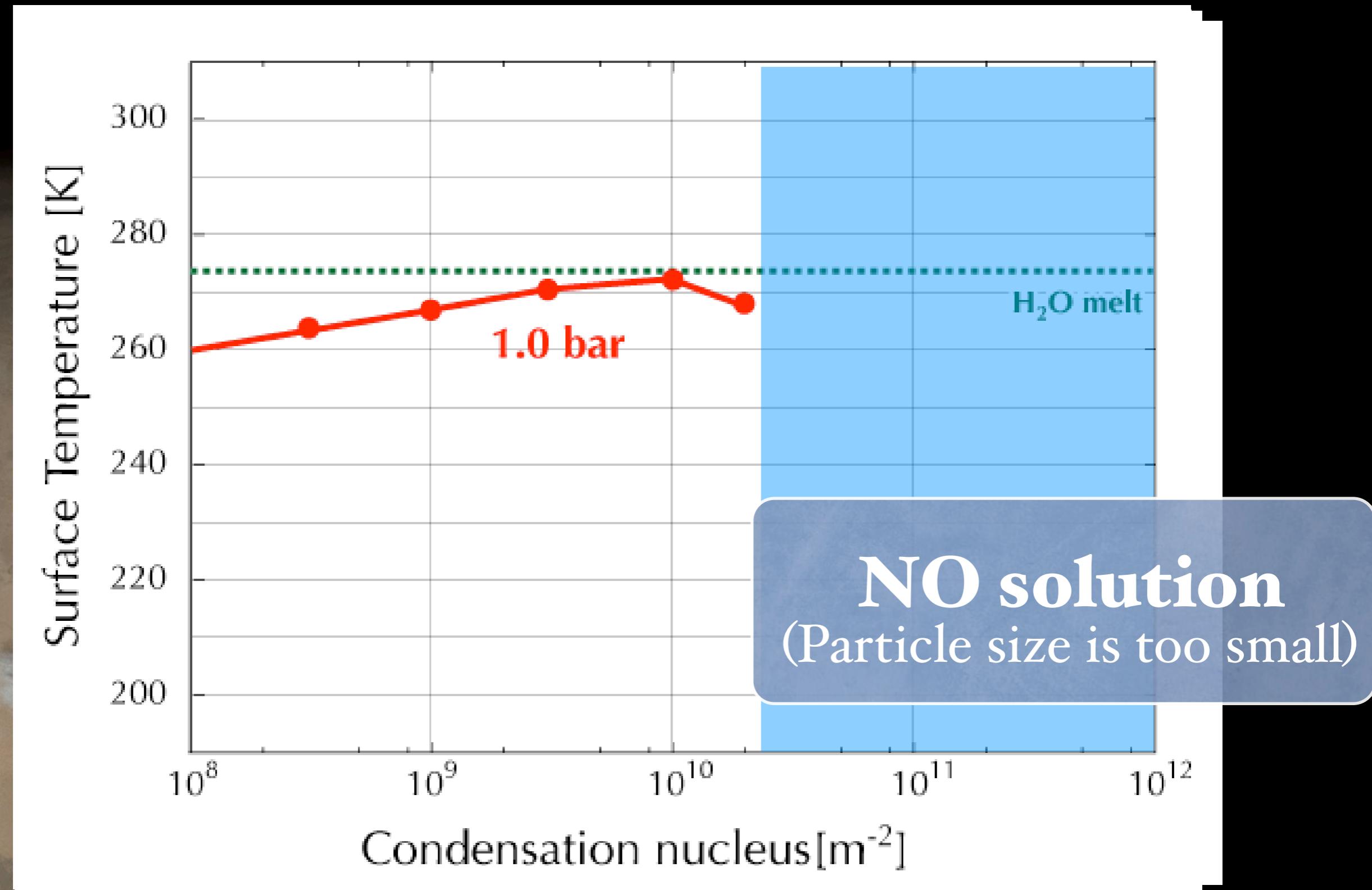
Surface temperature change  
to point C, in which these two  
equilibria are archived.



# Estimation of surface temperature

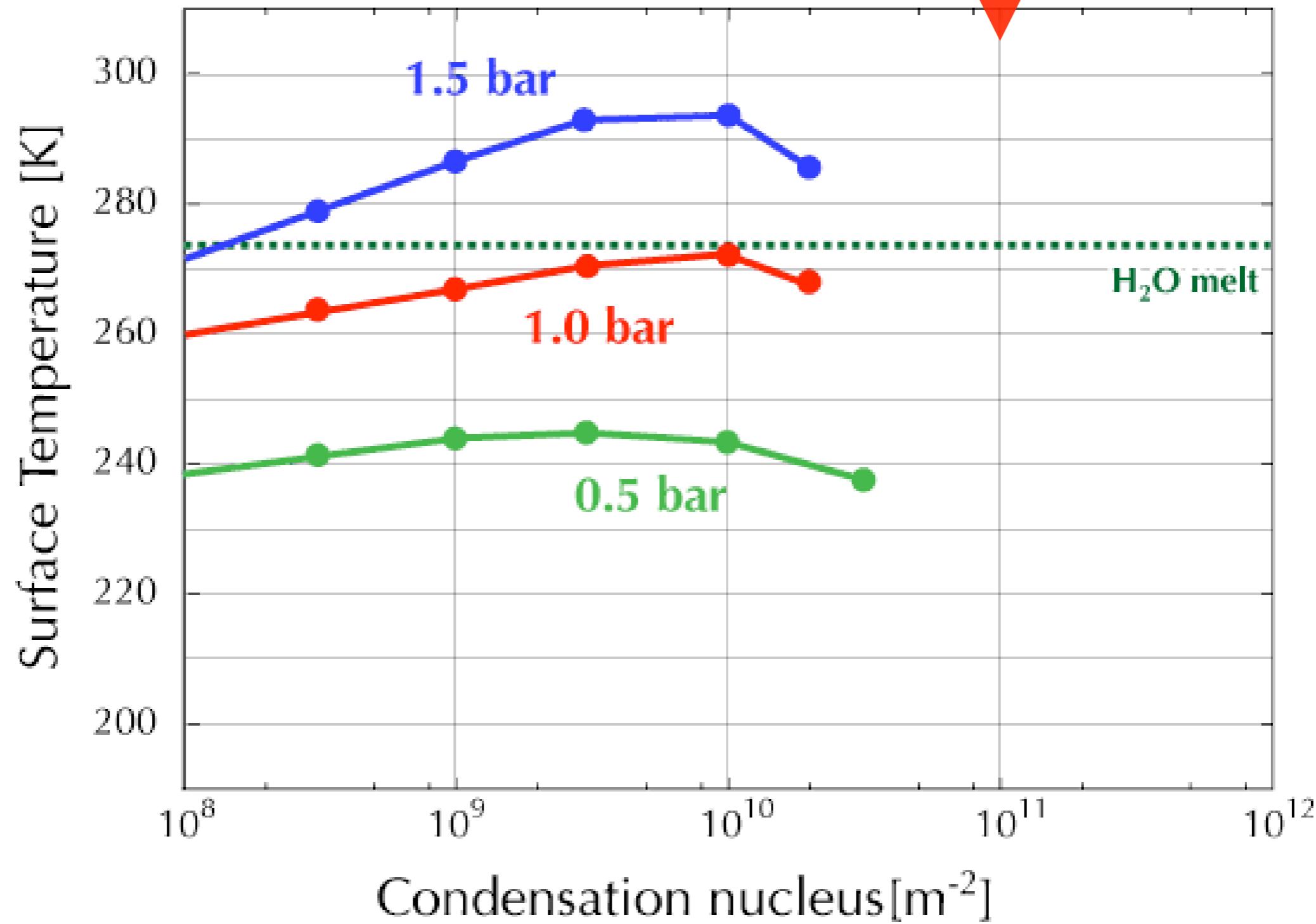


# Estimation of surface temperature



# Estimation of surface temperature

Dust (present Mars)



# Conclusion

- We construct simple 1-D radiation model and estimate surface temperature when atmospheric pressure and number of condensation nuclues are fixed.
  - the atmospheric pressure more than about 1 bar is necessary condition of the warm and wet climate on Early Mars.
- By the negative feedback of changes of the particle size and latent heat of condensation, the clouds may stabilize warm climate on early Mars.

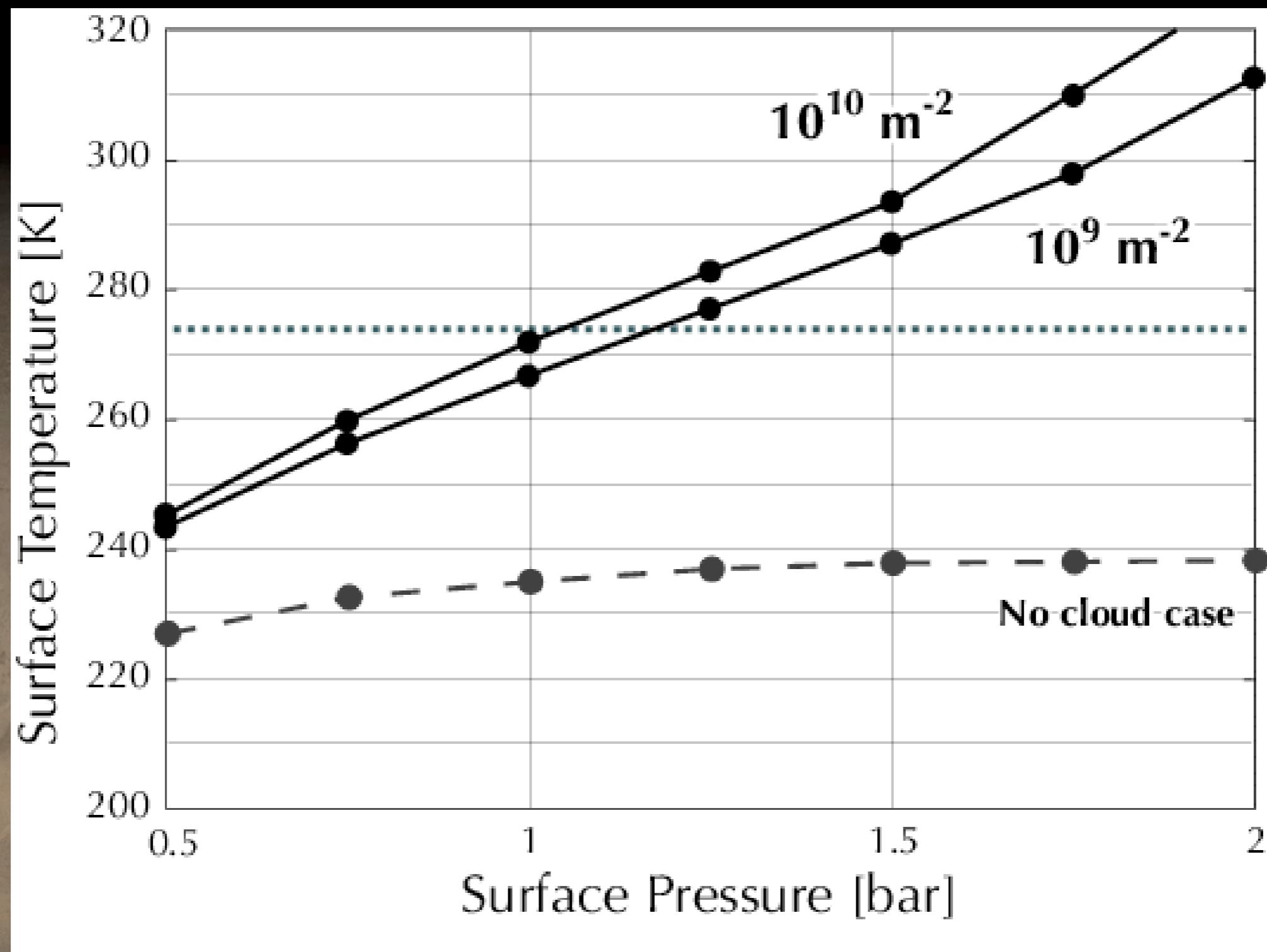
# References

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- Mischna, M. A., Kasting, J. F., and Freedman, R., 2000: Influence of carbon dioxide clouds on early Matrian climate, *Icarus*, **145**, pp.546–554
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# 質問, コメント

- 凝結核の見積もり方法としてアイデアは? 現在の地球の凝結核は、海の水しぶきが主である。過去の火星に海が存在していたならば、火星の凝結核もそうなることが想像されるので、地球の凝結核個数の比較はしてみましたか?(じょーじさん)
  - うーん、やってません... やってみます。
- メタンが存在している場合にはどうかわるのか?(Bullock さん)
  - メタンの温室効果も強まるが、おそらく対流圏の上層は放射冷却が効いて冷えるためにより厚い雲が生じるために雲による温室効果も強まり、より気候は暖まると考えている。
  - 要するにメタンが増えたら大気圧が増えた(一層あたりの透過率が下がった)ことと等価なわけで、地表面温度はあがるので?(じょーじさん)
- でもメタンが有ったら成層圏の温度があがるよ(by 大朝さん)
  - 実際に計算してたしかめてみます...(はて? なんでそうなるかは不明、メタンは太陽放射を吸収して地球のオゾンのような役割を果たすとか?? 赤外にしか吸収帯をもっていなければ放射冷却がより強く効くはずなのに...??)
- 凝結核が存在しない場合にも凝結は生じるはずでは?(今村さん)
  - 今回は凝結核がたくさん有る場合を考えているので、その効果は無視できる程度だと考えている。

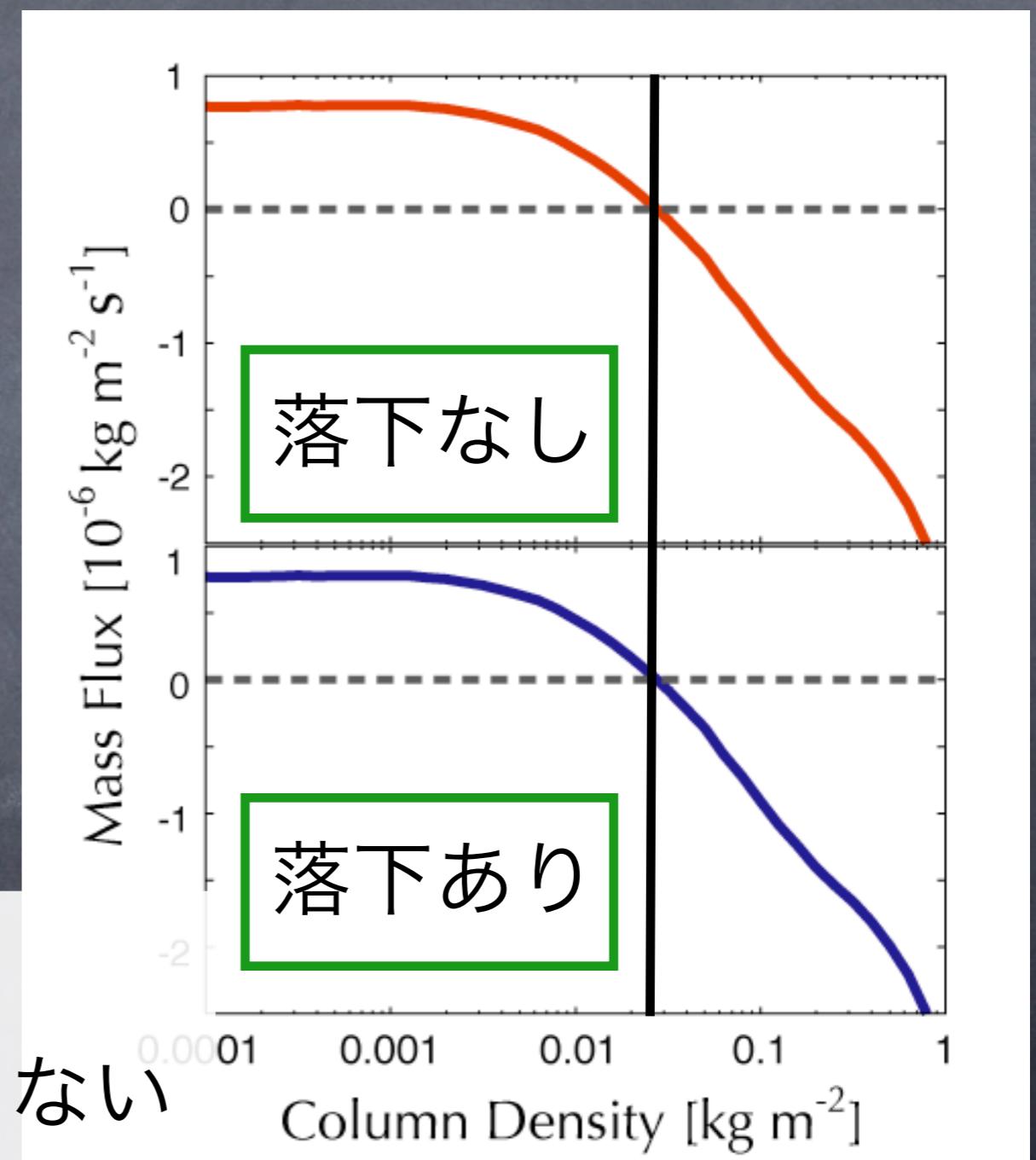
# 地表面温度の大気圧依存



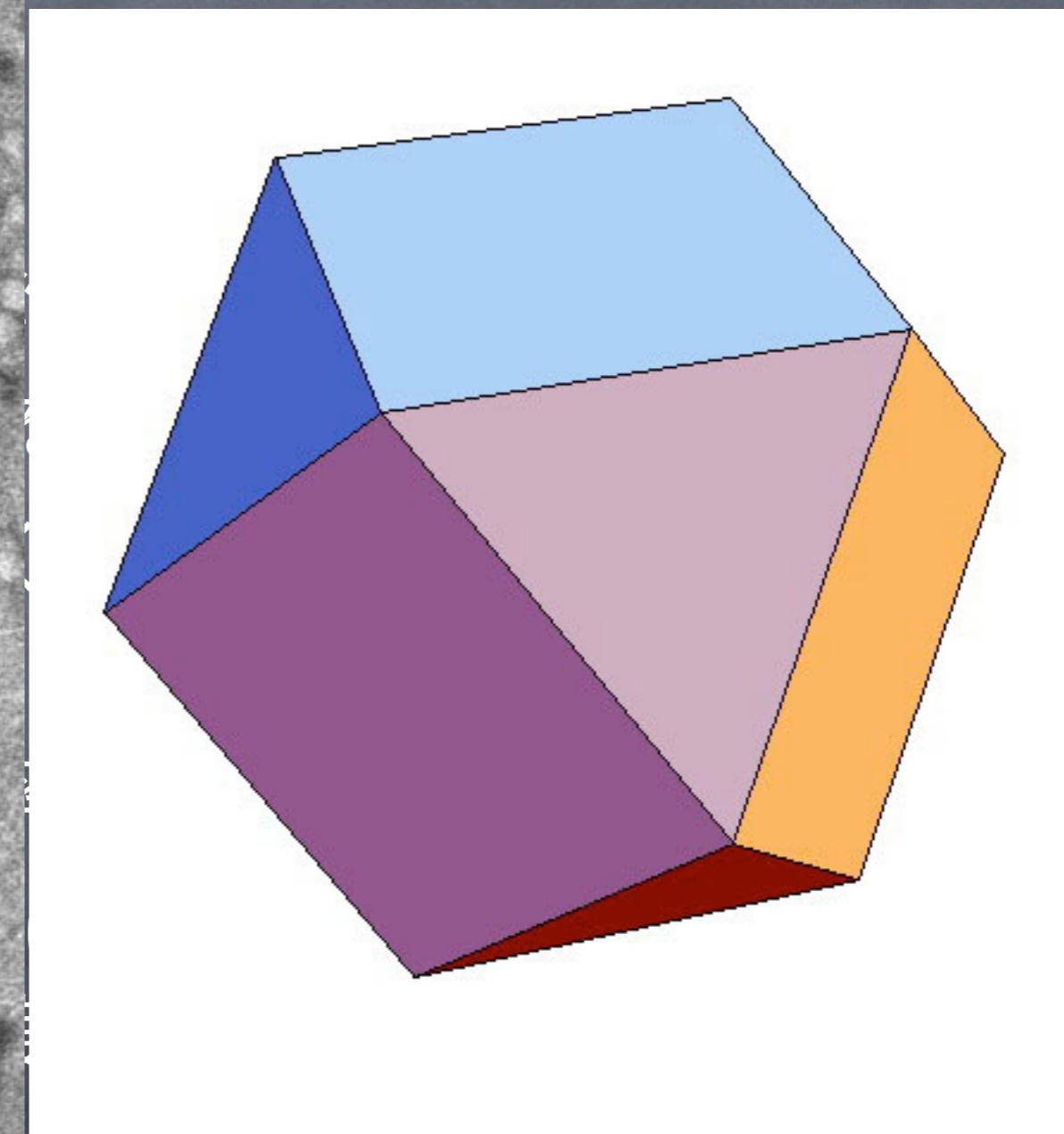
# 議論: 雲粒落下の影響

- 落下量の見積もり
- ストークス沈降速度
- 雲の鉛直構造は一様
- 雲層外に出た雲粒は蒸発

雲粒落下を考慮しても  
雲面密度の見積もりは変わらない

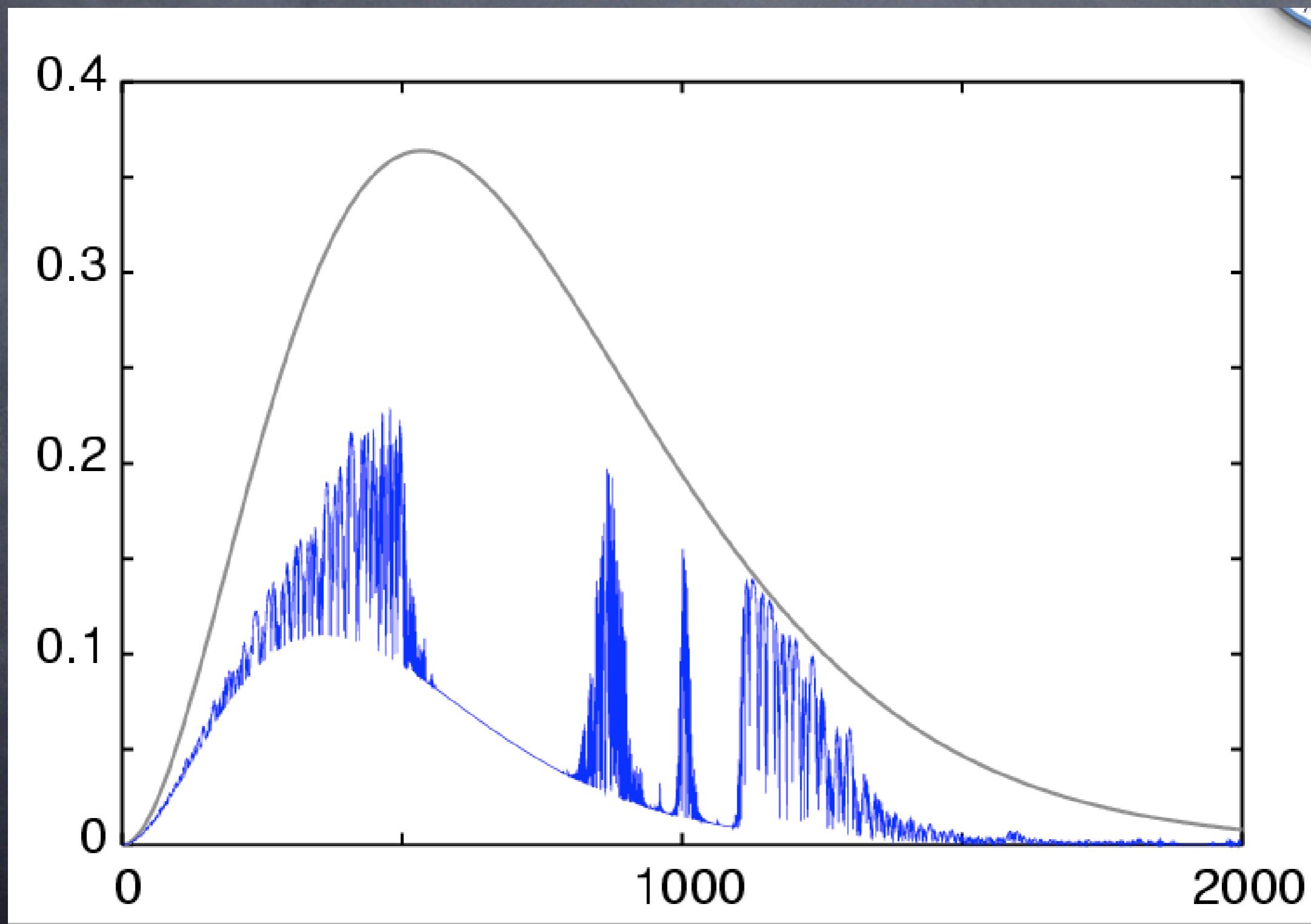


# 二酸化炭素の結晶構造



4snowflakes/cuboctahedronrh400.jpeg

# 雲への入射スペクトル



# 地球型惑星の比較

	火星	地球	地球 (堆積岩)	金星
大気組成				
N2	2.7	78.1	1.0	1.8
O2	-	20.9	-	-
Ar	1.6	0.9	0.01	0.02
CO2	95.3	0.035	99.0	98.1
CO2 分圧 [bar]	$6 \times 10^{-3}$	$10^{-4}$	80	90